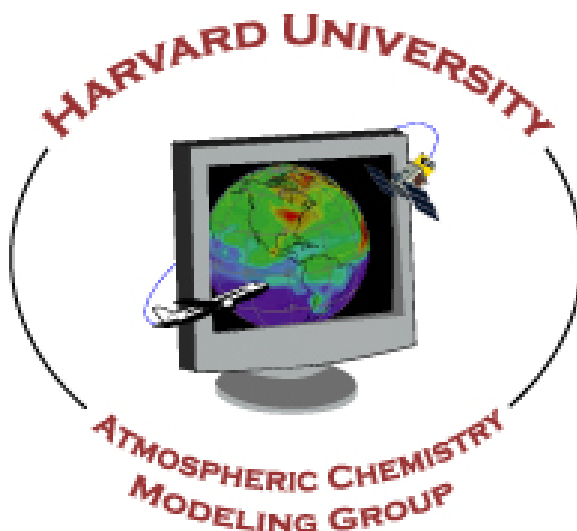


# **Sensitivity of sulfate direct climate forcing to the particle physical state: A global perspective**

**Jun Wang, Andrew Hoffman, Scot Martin, Daniel Jacob**



Present at AEROCENTER/GSFC, Nov. 21, 2006

# Sulfate physical state and chemical composition

Neutralization  $X = [\text{NH}_4]/2[\text{SO}_4]$

## Solids:

AS	$(\text{NH}_4)_2\text{SO}_4$	$X = 1$
LET	$(\text{NH}_4)_3\text{H}(\text{SO}_4)_2$	$X = 0.75$
AHS	$(\text{NH}_4) \text{HSO}_4$	$X = 0.5$

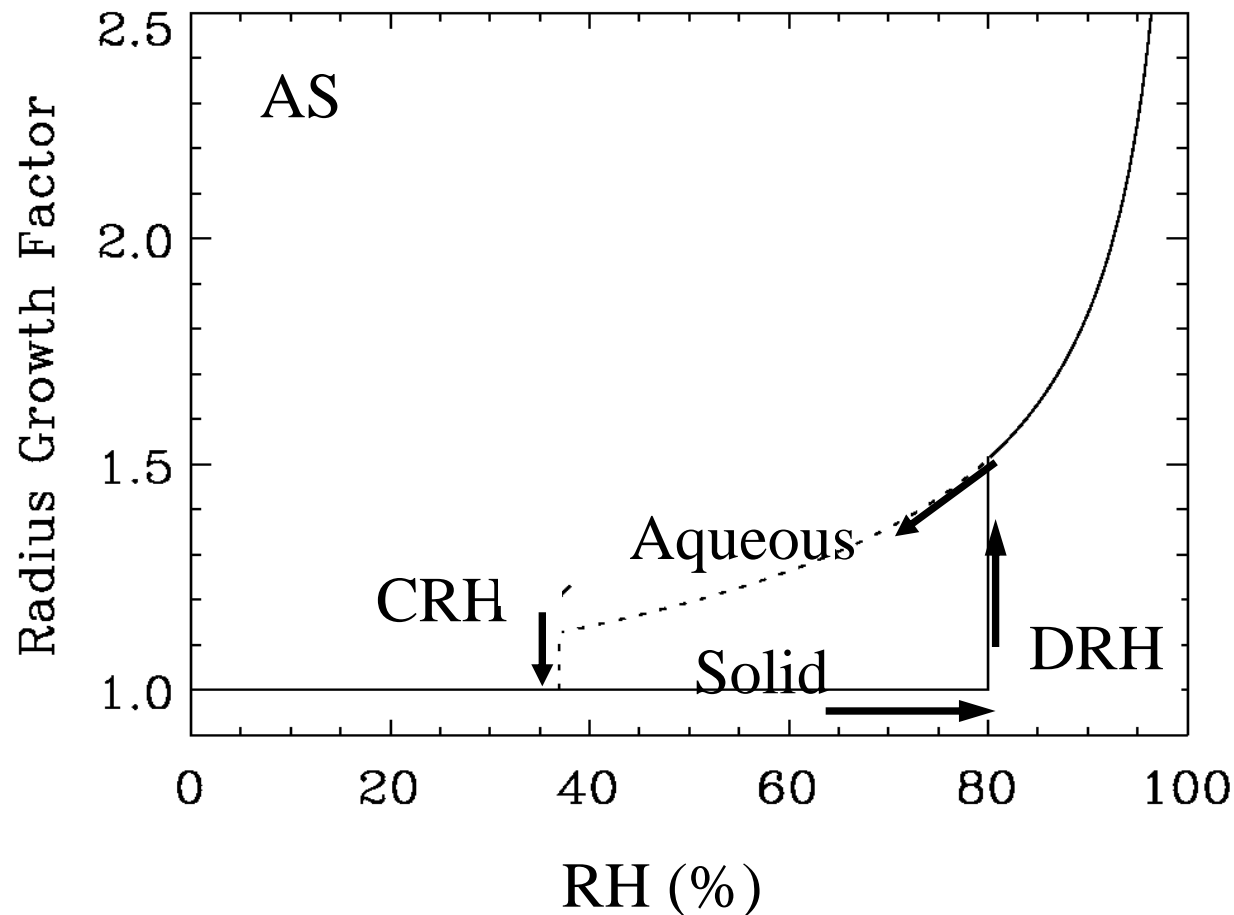
## Aqueous:

SA	$\text{H}_2\text{SO}_4, \text{H}_2\text{O}$	$X = 0$
SO4aq	$\text{SO}_4^{2-}, \text{H}^+, \text{NH}_4^+, \text{H}_2\text{O}$	$0 < X < 1$

# Hysteresis Loop of Sulfate Hygroscopicity

CRH: Crystalline relative humidity

DRH: Deliquesce relative humidity



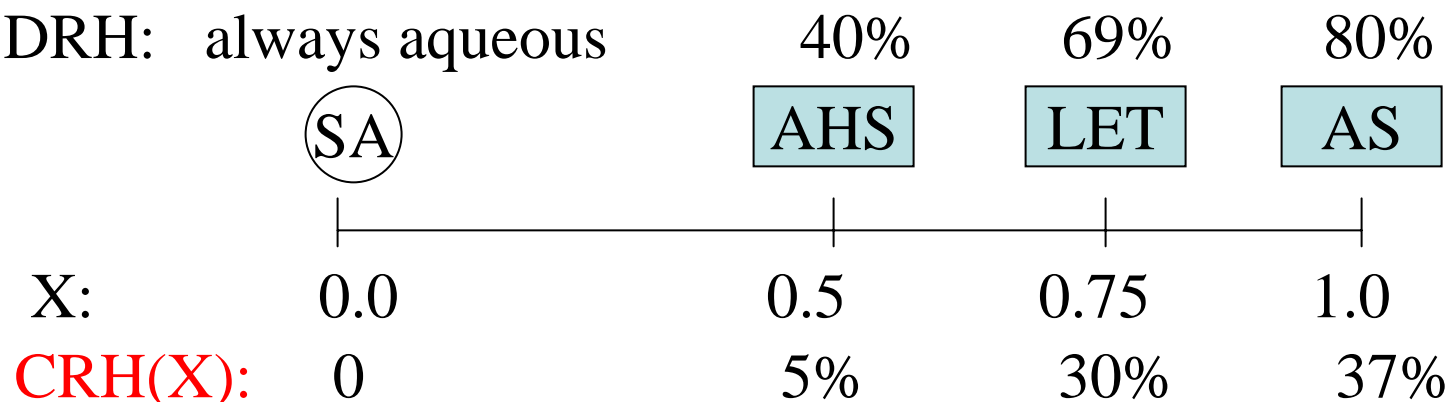
3 factors:

- (1) local RH
- (2) RH history
- (3) CRH & DRH

$CRH < RH < DRH$   
solids or aqueous?

# Composition-dependence of CRH and DRH

DRH: always aqueous



Polynomials of X  
*Martin et al.*, 2003

An example:

aqueous particles with  $X = 0.9$  ( $\text{CRH}(x) = 32\%$ ) at  $\text{RH} = 85\%$

- 1) Decreasing RH to 70%, all aqueous
- 2) Decreasing RH to 32%, all solid particles (LET&AS).
- 3) Increasing RH to 70%, mixed phase (aqueous LET and solid AS)

# Importance of Sulfate Physical State

- Microphysical importance
  - Particulate Matter (PM) Air quality (particle size and mass)
  - Heterogeneous chemistry (hydrolysis of  $\text{N}_2\text{O}_5$ )
  - Cloud formation (ice/water CCN)
- Radiative importance
  - Aerosol refractive index and size
  - Sulfate direct climate forcing (SDCF)
  - Visibility

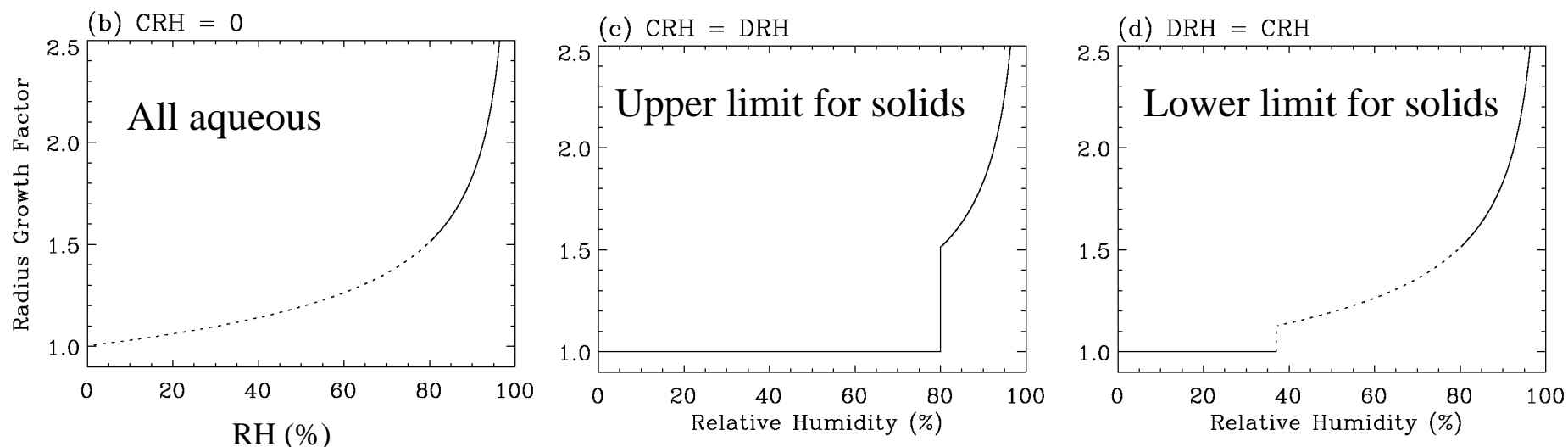
# Previous Studies on Sulfate Phase Transition

CTM:

- 1) No phase, only  $\text{SO}_4$  mixing ratio
- 2) Diagnosis phase based on local  $X$  and  $\text{RH}$ , with assumed  $\text{RH}$  history.
- 3) track  $\text{RH}$  history using Lagrangian model, and diagnosis the phase.

Radiative calculation:

- 1) CRH and DRH equal to a  $\text{RH}$  threshold to remove bifurcation.
- 2) Other parameterization methods



# An example

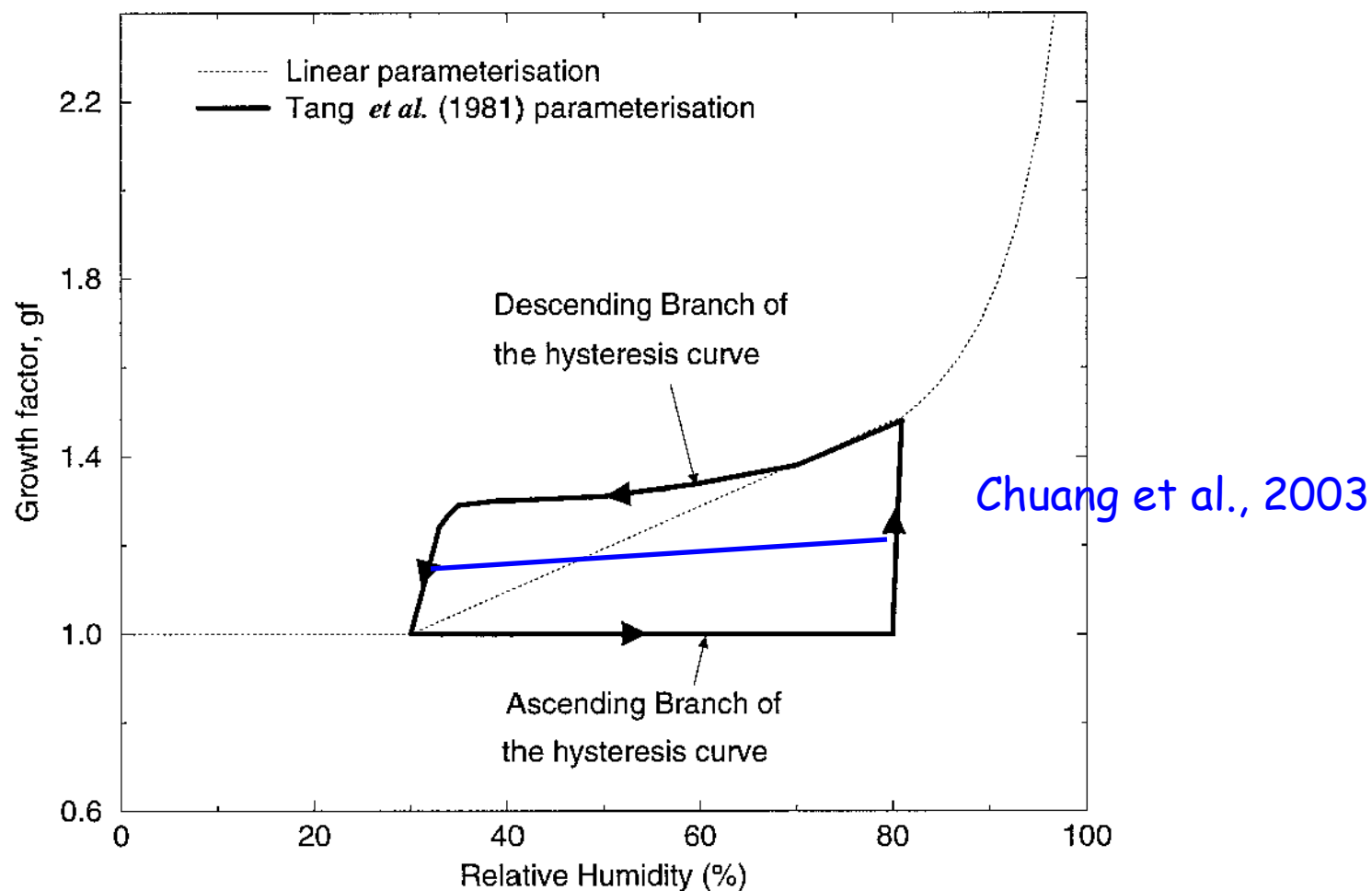


FIG. 1. The growth factor for ammonium sulfate aerosol as a function of relative humidity for the linear, ascending, and descending relative humidity schemes described in the text.

Haywood *et al.*, 1997  
with modification

# Uncertainty Envelope from Previous Studies

Assuming CRH and DRH of AS for all sulfate particles

Aerosol type	Levels	Relative humidity parameterization	Cloud scheme	NH W m <sup>-2</sup>	SH W m <sup>-2</sup>	Global W m <sup>-2</sup>	Global change %
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	7	linear	UM	-0.60	-0.15	-0.38	0
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	7	linear	cloud mask	-0.59	-0.14	-0.36	-5
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	5	linear	UM	-0.67	-0.17	-0.42	+11
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	10	linear	UM	-0.57	-0.14	-0.36	-5
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	7	dry	UM	-0.38	-0.09	-0.24	-37
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	7	descending	UM	-0.63	-0.16	-0.39	+3
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub>	7	ascending	UM	-0.54	-0.14	-0.34	-11
H <sub>2</sub> SO <sub>4</sub>	7	d'Almeida et al. (1991)	UM	-0.66	-0.17	-0.41	+8

Haywood et al., 1997

Using CRH and DRH from aerosol thermodynamical model.

( $F_L$  and  $F_U$ , respectively, in W m<sup>-2</sup>) are calculated. Including both anthropogenic and natural emissions, we obtain global annual averages of  $F_L = -0.750$ ,  $F_U = -0.930$ , and

$\Delta F_{U,L} = 24\%$  for full sky calculations without clouds and  $F_L = -0.485$ ,  $F_U = -0.605$ , and  $\Delta F_{U,L} = 25\%$  when clouds are included. Regionally,  $\Delta F_{U,L} = 48\%$  over the USA, 55%

Martin et al, 2004.



# Questions

- 1) Mass percentage of solid sulfate?
- 2) Contribution of solids to sulfate direct climate forcing (DCF) ?

A systematic bias or error (not  $\pm$  random uncertainty) !

# Approach

- 1) A box model approach
- 2) GEOS-CHEM investigation

# Box Model Estimate

Under thin-layer approximation (*Wiscombe and Grams, 1976*):

$$SDCF = -A \left( \underbrace{\omega_{sd} \bar{\beta}_{sd} \tau_{sd}}_{\text{Solid}} + \underbrace{\omega_{aq} \bar{\beta}_{aq} \tau_{aq}}_{\text{Aqueous}} \right)$$

$$A = -\frac{1}{2} S_0 T (1 - A_c)(1 - R_s)^2$$

$S_0$ : solar constant

$A_c$ : cloud fraction

$R_s$ : surface reflectance

$T$ : atmos. transmittance

$\omega$ : single scattering albedo (= 1)

$\bar{\beta}$ : time-averaged backscattering fraction

$\tau$ : aerosol optical thickness

Climate System parameters
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Forcing Agent parameters
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# Importance of Solids on Forcing Estimate

Aerosol term:  $\bar{\beta}_{sd} \tau_{sd} + \bar{\beta}_{aq} \tau_{aq}$

$$= \bar{\beta}_{sd} B_{sd} E_{sd} + \bar{\beta}_{aq} B_{aq} E_{aq}$$

Where

E: mass extinction efficiency  $\text{m}^2 (\text{gSO}_4^{2-})^{-1}$

B: sulfate burden  $(\text{gSO}_4^{2-}) \text{m}^{-2}$

$B_{sd}/B_{aq}$  is unknown, previous studies use:

$$= \bar{\beta}' E' B$$

$$= \bar{\beta}' E_{sd} G_{\tau} B$$

where

$$\bar{\beta}' E' = \frac{\bar{\beta}_{sd} E_{sd} B_{sd} + \bar{\beta}_{aq} E_{aq} B_{aq}}{B}$$

$$B = B_{sd} + B_{aq}$$

Where

$$G_{\tau} = G_E - \frac{B_{sd}}{B} (G_E - 1)$$

$$G_E = \frac{E_{aq}}{E_{sd}}$$

Previous common assumption:

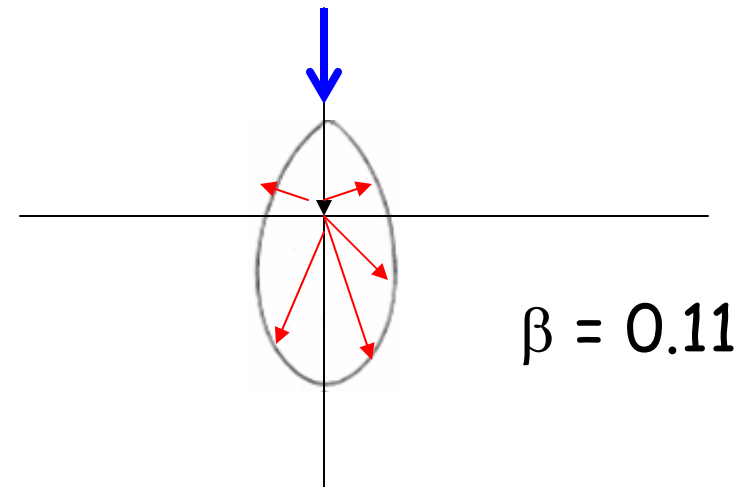
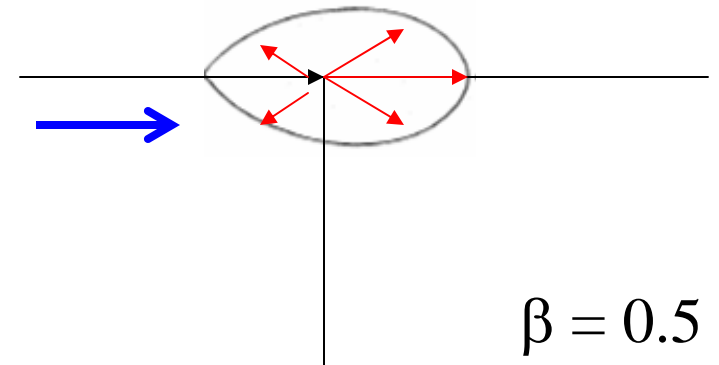
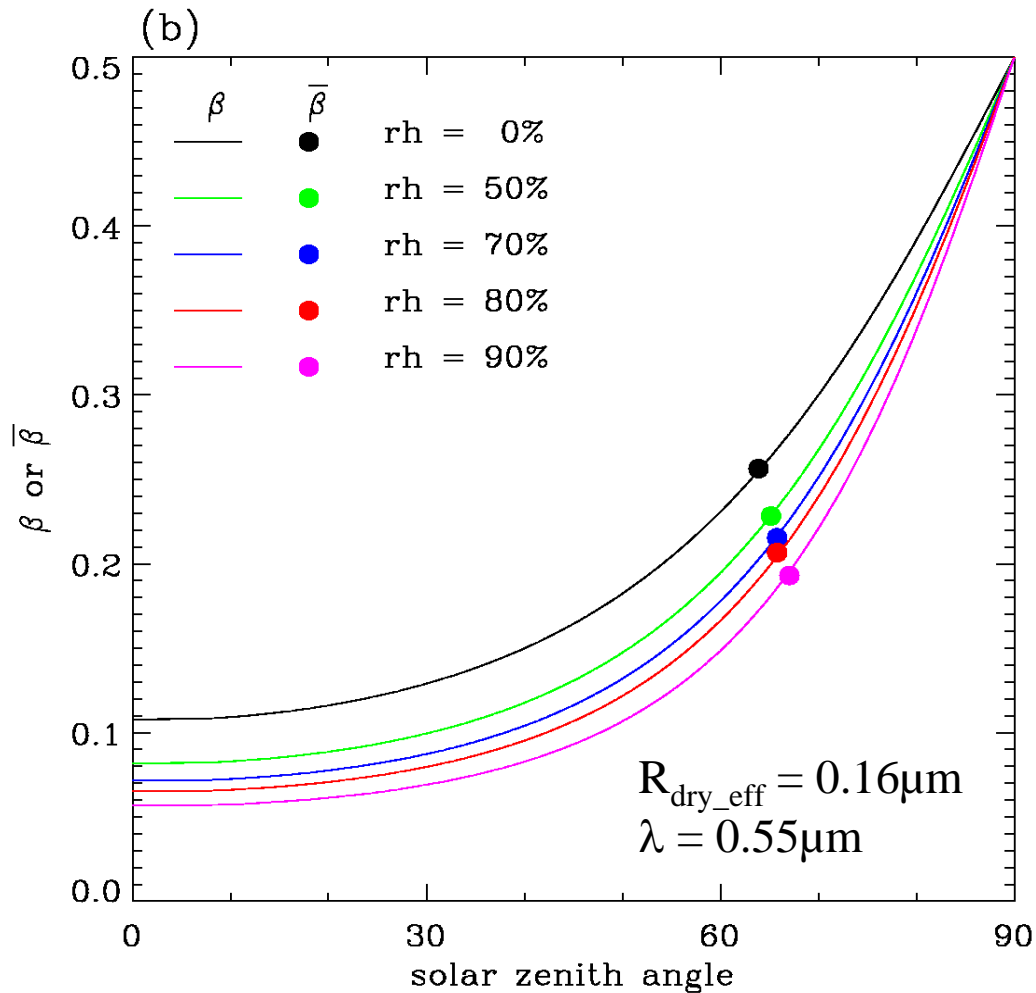
$$G_{\tau} = G_E$$

$$\bar{\beta}' = \bar{\beta}_{sd} \quad \text{or} \quad \bar{\beta}' = \bar{\beta}_{aq}$$

$G_{\tau}$ : Optical thickness growth factor

$G_E$ : Mass extinction efficiency growth factor

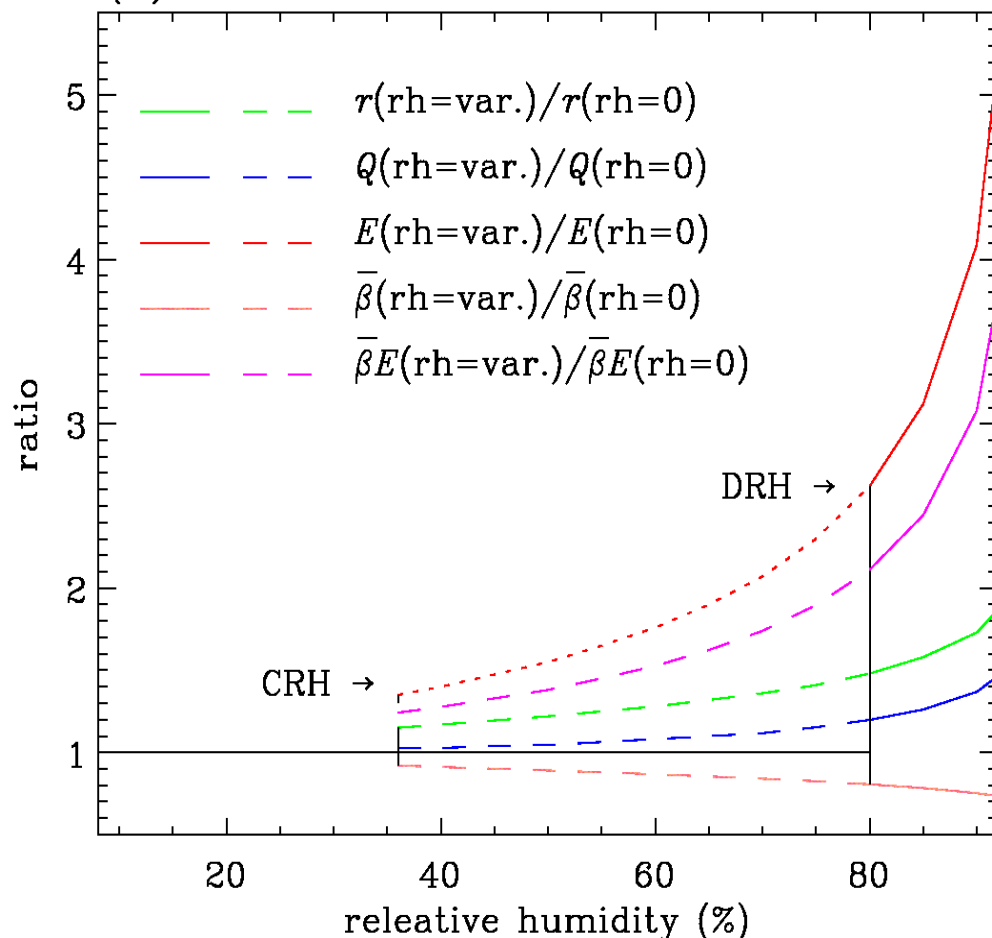
# $\overline{\beta}$ : time-averaged backscattering fraction



At RH = 80%,  $\overline{\beta}_{aq} = 0.7 \overline{\beta}_{sd}$

$$\overline{\beta} = \int_{\mu_{\min}}^{\mu_{\max}} \beta(\mu) d\mu$$

# sulfate physical state impact on scattering properties



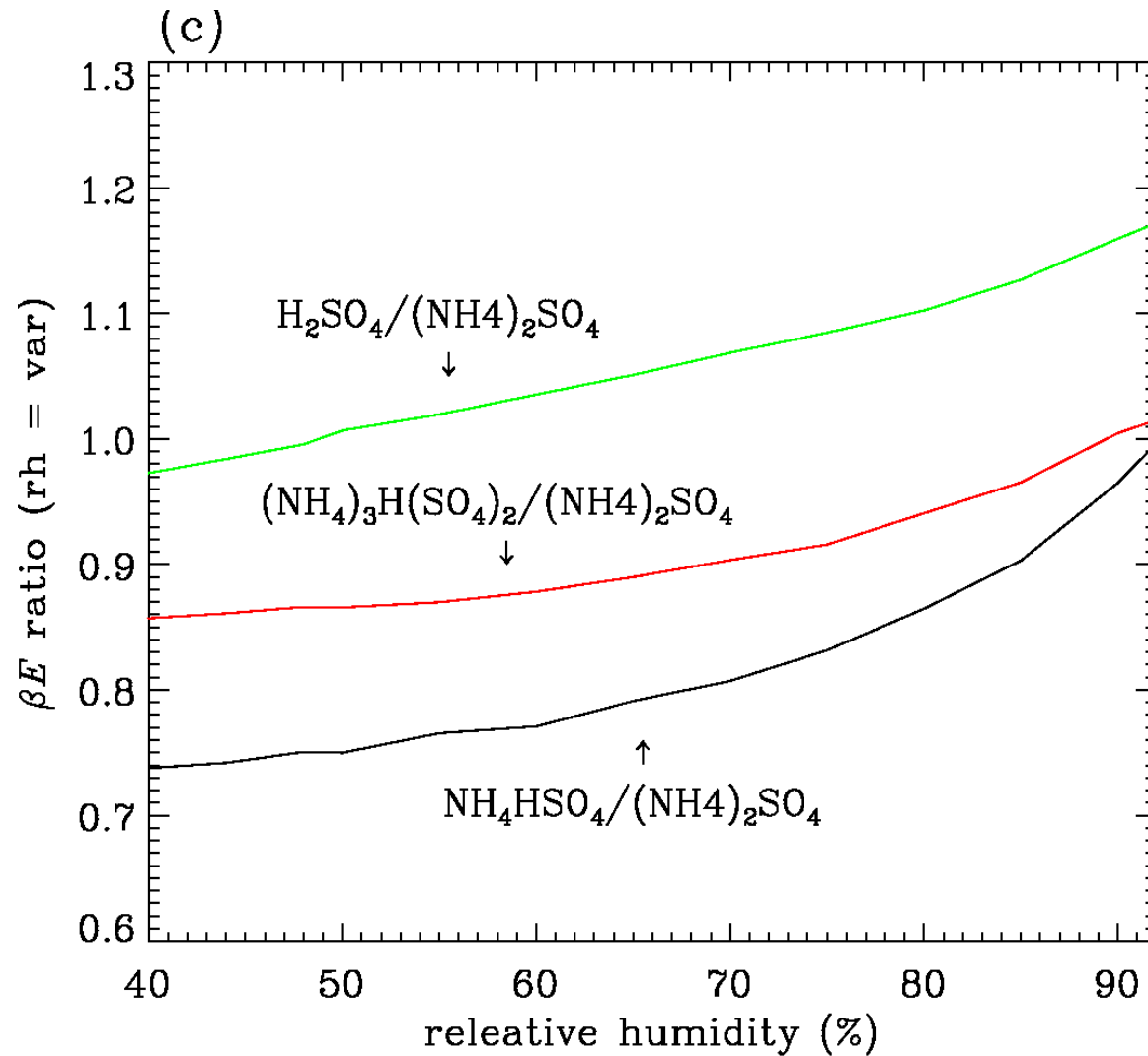
For aqueous particles, as RH increases,

$r$  increases,  
 $Q$  (extinction efficiency)  
and  $E$  increase

However,  $\beta$  decreases

Overall, increase the aerosol forcing.

# Impact of Chemical Composition



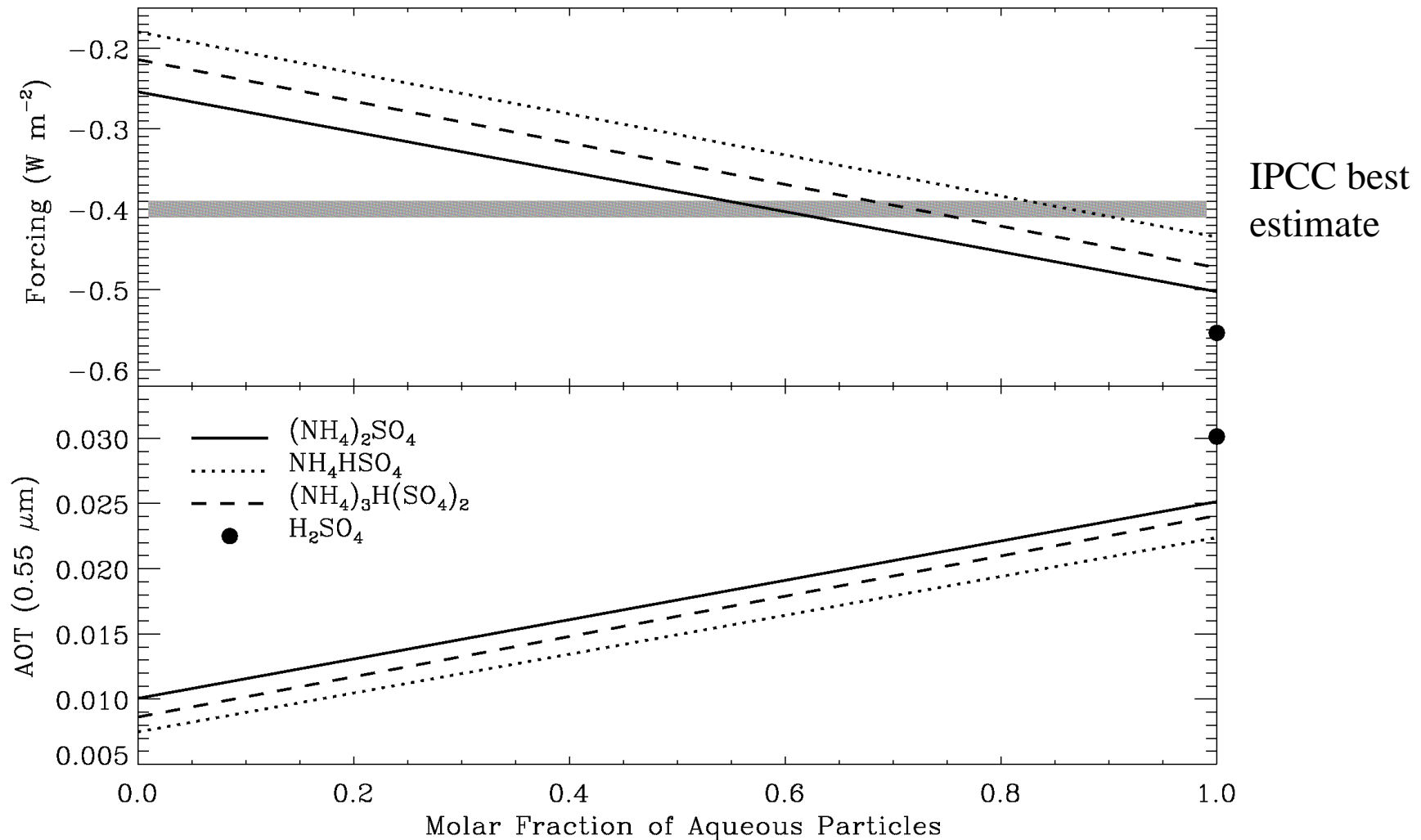
$$\overline{\beta}_{aq} \tau_{aq}$$

20% difference due to  $\Delta$ composition

$$\overline{\beta}_{sd} \tau_{sd}$$

30% difference due to  $\Delta$ composition

# Sulfate AOT and forcing for variable $B_{sd}/B$



The optical properties for aqueous particles at RH=80% is used. Climate system parameters are same as Charlson et al (1992)

# GEOS-CHEM Investigation

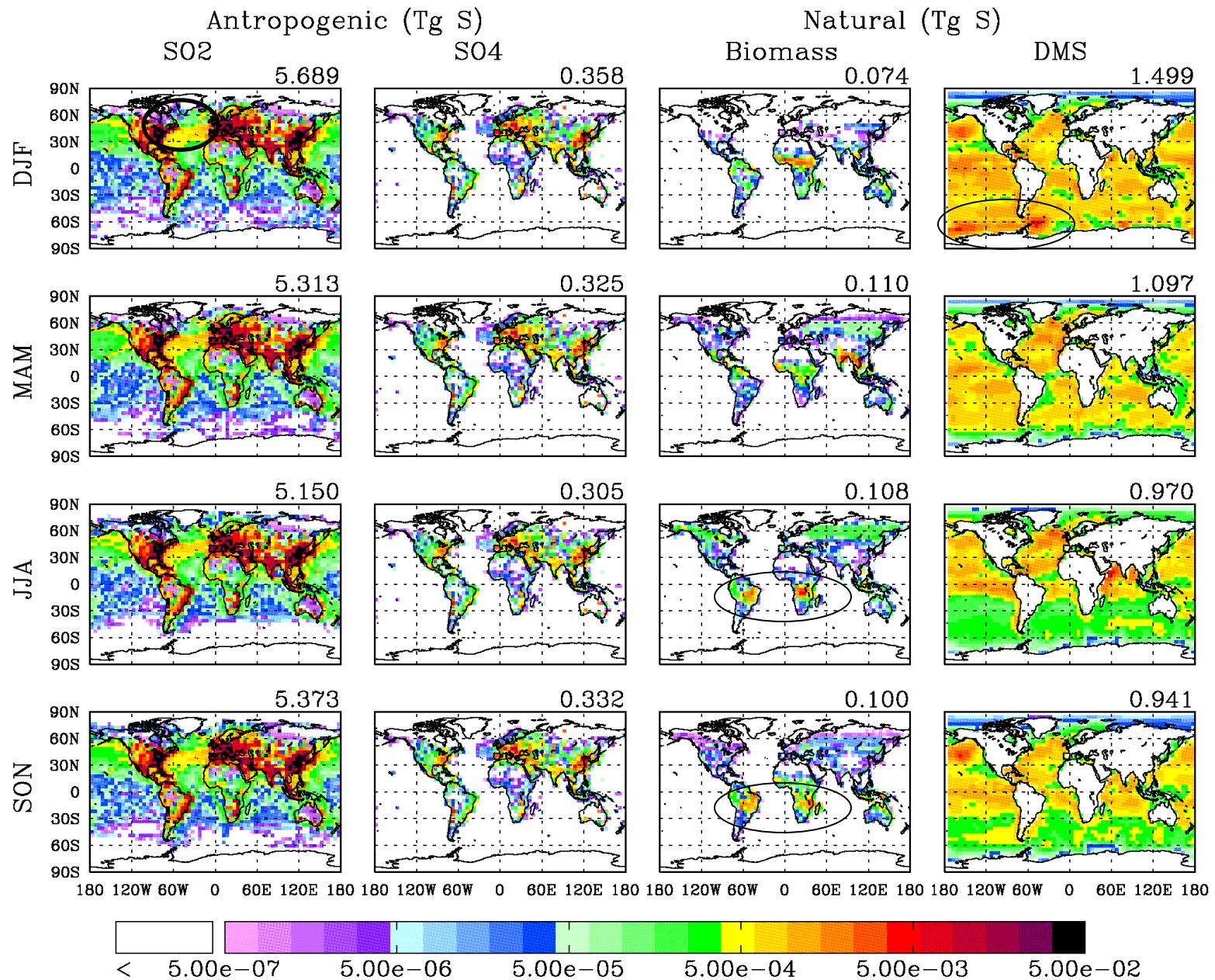
- 1) Seasonal and geographical distribution of  $B_{sd}$
- 2) Percentage of solid particle contribution to the global full-sky anthropogenic forcing
- 3) Forcing difference in the following 4 cases, basecase, CRH = 0, CRH=DRH, and DRH=CRH.
- 4) Where and how much would be largest forcing difference caused by the sulfate phase transition (regional perspective).



# Model development

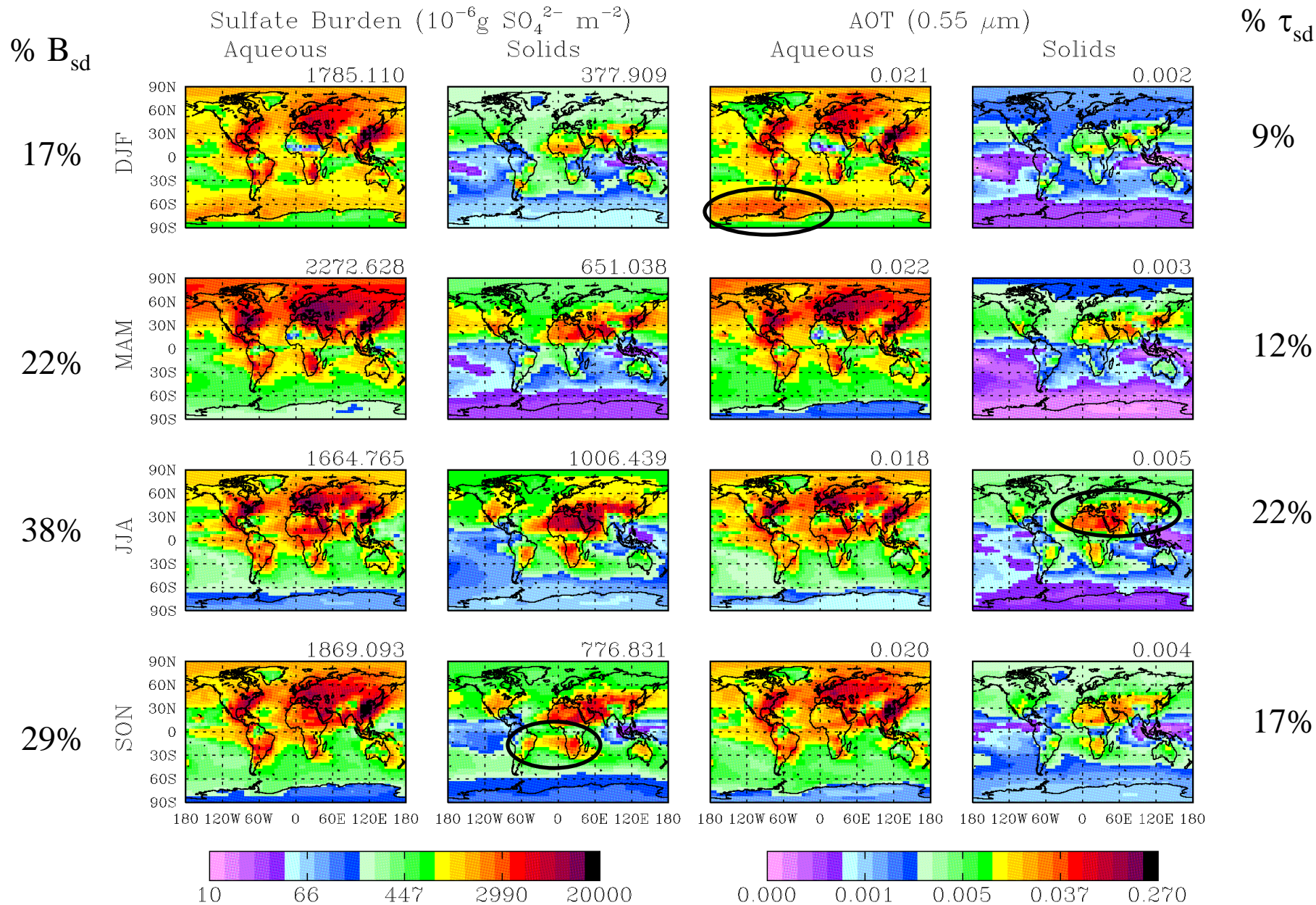
- Based upon Park et al (2004), with the following modification:
  - No nitrate.  $\text{H}_2\text{SO}_4\text{-H}_2\text{O}$  system only.
  - Sulfate mass are transported in 4 species: solid AS, LET, and AHS particles, aqueous  $\text{SO}_4$ .
  - CRH (X) and DRH of solids from Martin et al., 2003.
  - The concentration of each sulfate species are calculated according to the CRH(X), X, ambient RH, and DRH values.
- Because we explicitly track the solid and aqueous phase at each time step and model grid, the RH history on sulfate phase is retained.

# Emissions: ~80% are anthropogenic



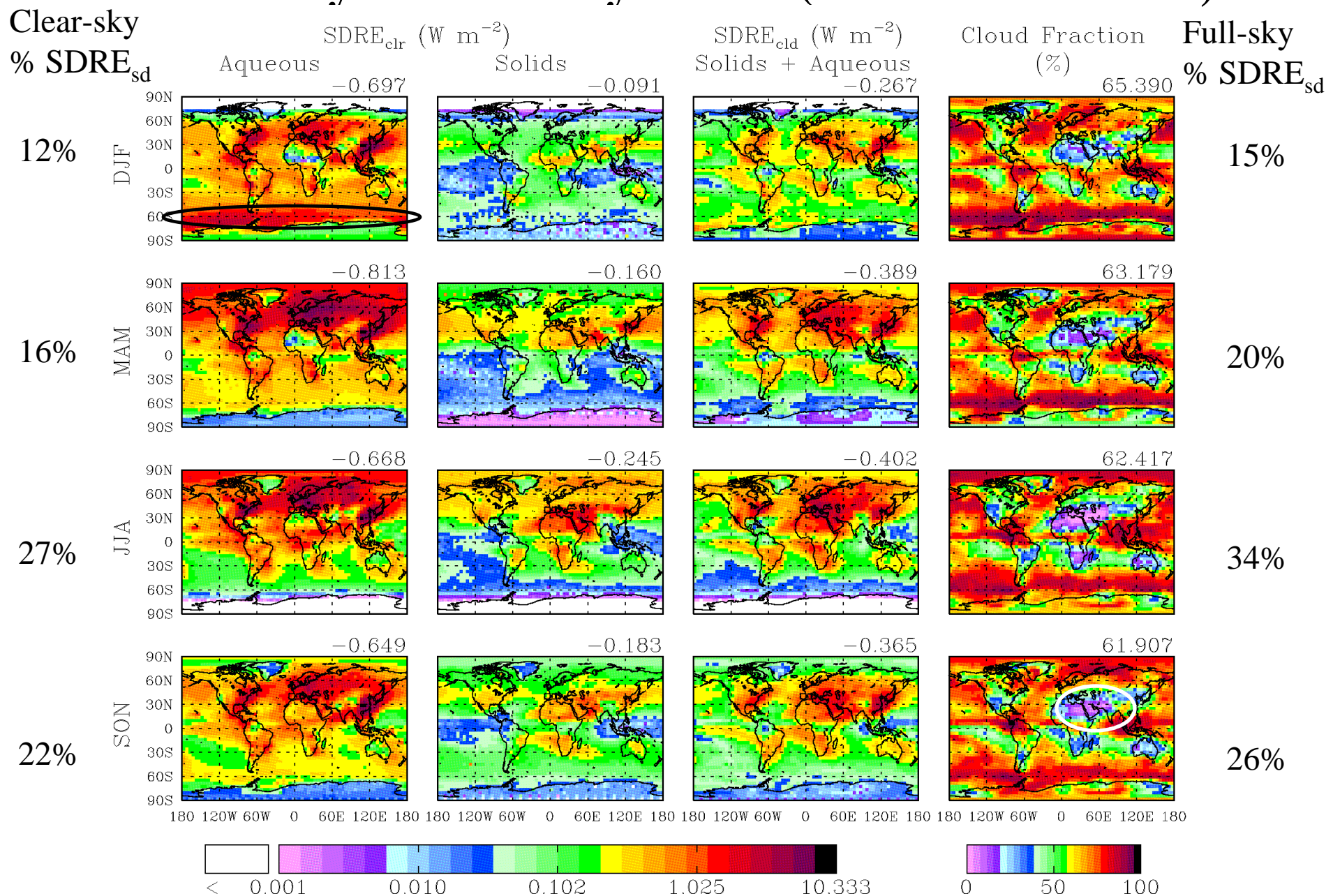
Following Park et al (2004), we consider biomass burning as natural emission.

# Seasonal $\text{SO}_4^{2-}$ burden and AOT in base case



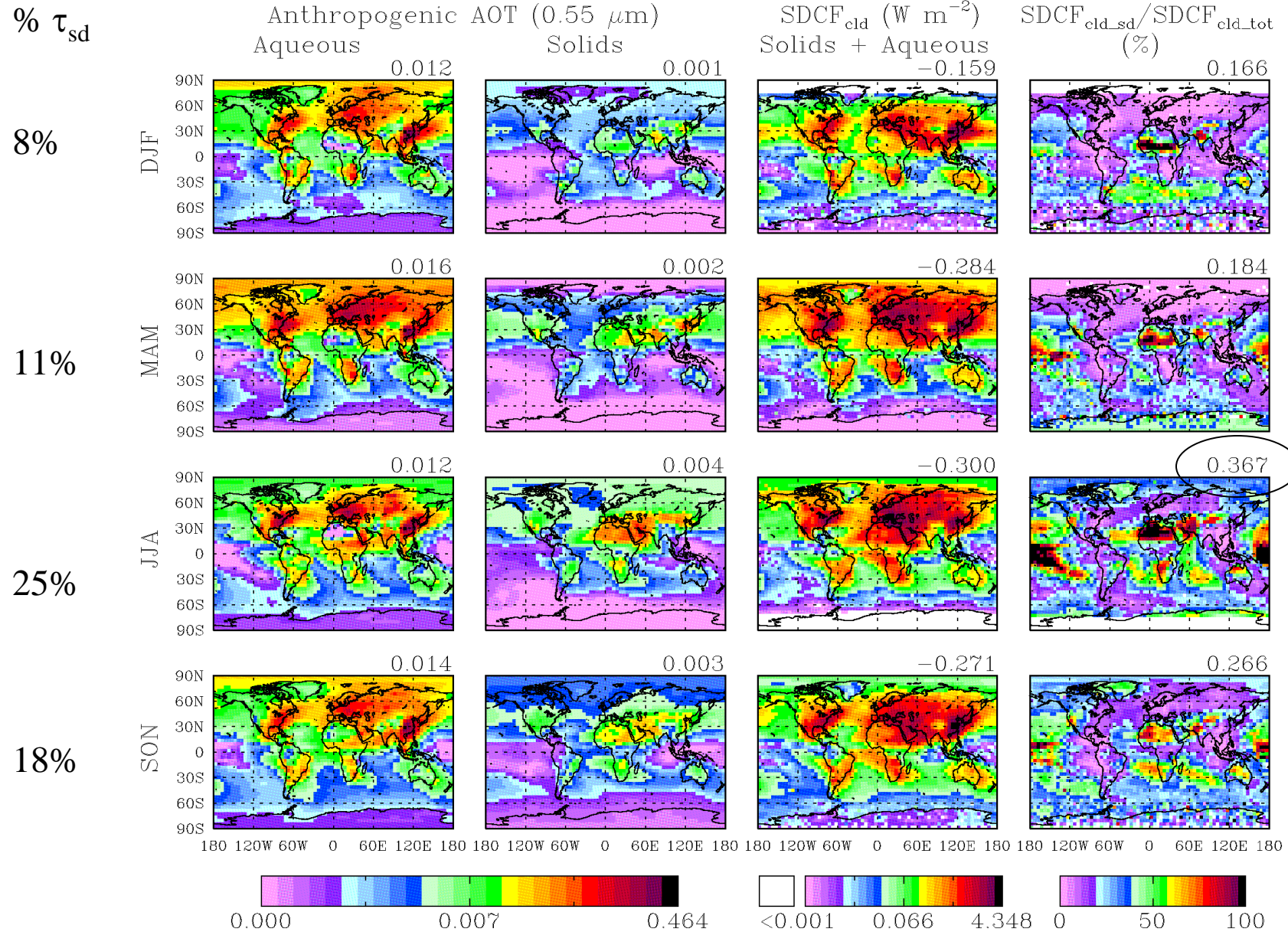


# Clear-sky and full-sky SDRE (natural + anthro.)



Covariance matters. Cloud distribution favors the forcing of solids.

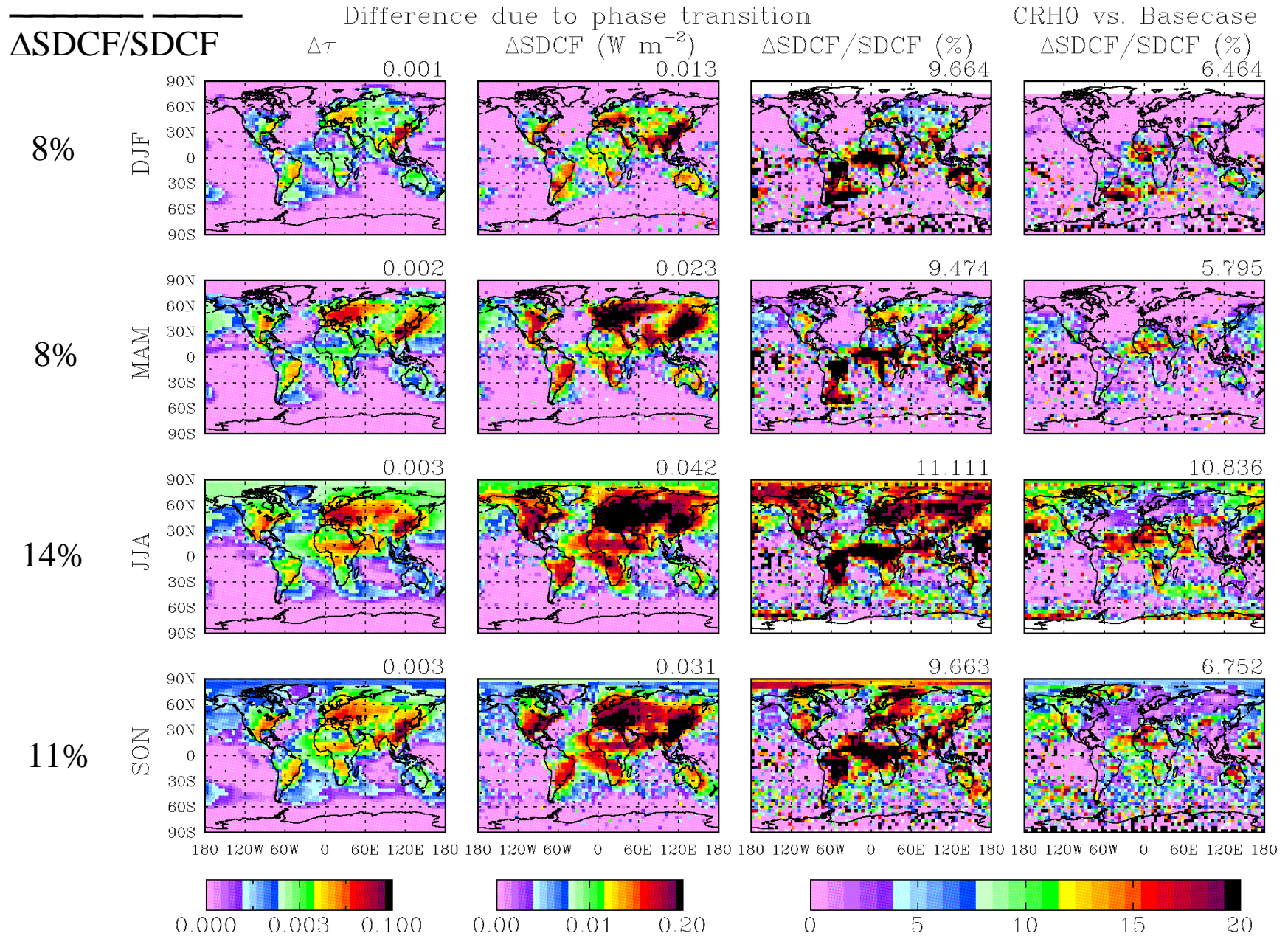
# SDRF (~70% of SDRE)



Forcing difference is over the Sahel region.



# Effect of particle state on SDCF



# Quantitative Summary of SDRE

	Basecase	CRH=0	CRH=DRH	DRH=CRH
<b>B (mg SO<sub>4</sub> m<sup>-2</sup>)</b>				
Total	2.601	2.591	2.609	2.590
% due to solids	27.0%	0.0%	48.3%	12.7%
<b><math>\tau</math> at 0.55<math>\mu</math>m (*10000)</b>				
Total	240	253	217	247
% due to solids	15.4%	0.0%	30.0%	6.9%
<b>SDRE<sub>cld</sub> (wm<sup>-2</sup>)</b>				
Total	0.356	0.371	0.332	0.363
% due to solids	24.4%	0.0%	44.0%	11.8%
<b>% to Basecase</b>				
$\tau$	100%	105.4%	90.4%	102.9%
<b>SDRE<sub>cld</sub></b>	<b>100%</b>	<b>105.1%</b>	<b>93.0%</b>	<b>101.9%</b>

# Quantitative Summary of SDCF

	basecase	CRH=0	CRH=DRH	DRH=CRH
<b>B (mg SO<sub>4</sub> m<sup>-2</sup>)</b>				
Total	1.859	1.849	1.866	1.851
Solids %	28.1%	0%	52.4%	13.0%
<b><math>\tau</math> at 0.55<math>\mu</math>m (*10000)</b>				
Total	166	175	136	171
Solids %	16.9%	0%	36.8%	7.0%
<b>SDCF (wm<sup>-2</sup>)</b>				
Total	0.254	0.266	0.240	0.259
Solids %	25.6%	0%	47.5%	12.0%
<b>Compared to base case</b>				
$\tau$ % to basecase	100%	105.4%	81.9%	103.0%
<b>SDCF<sub>cld</sub> %</b>	<b>100%</b>	<b>104.7%</b>	<b>94.5%</b>	<b>102.0%</b>

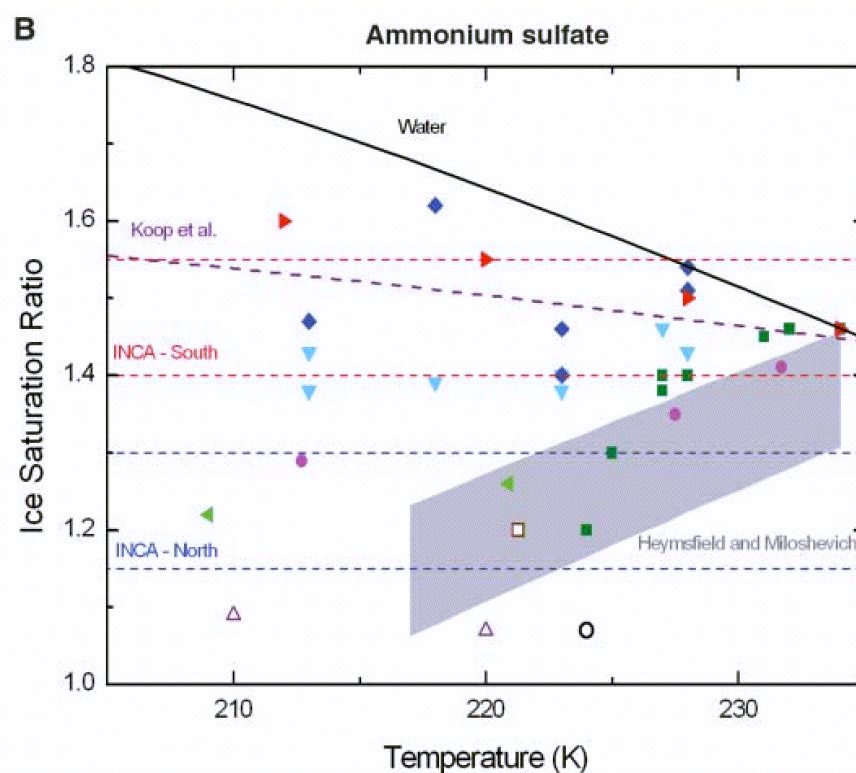
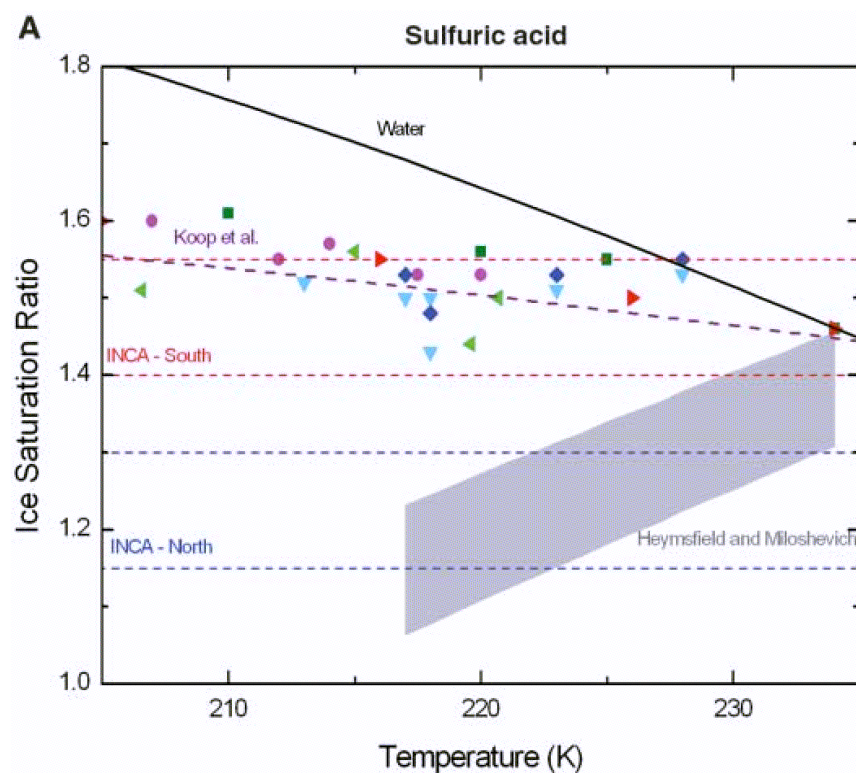


# Discussion

## Solid Ammonium Sulfate Aerosols as Ice Nuclei: A Pathway for Cirrus Cloud Formation

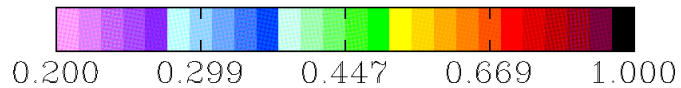
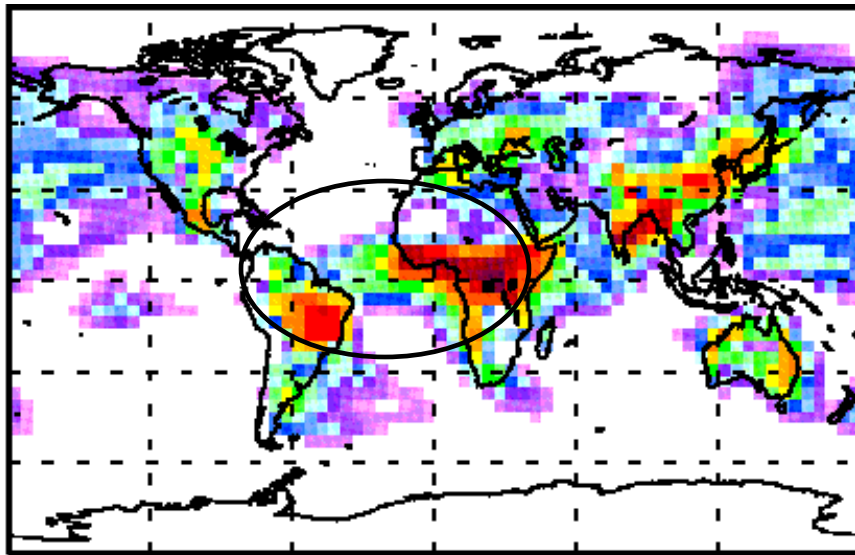
J. P. D. Abbatt,<sup>1\*</sup> S. Benz,<sup>2</sup> D. J. Cziczo,<sup>3</sup> Z. Kanji,<sup>1</sup> U. Lohmann,<sup>3</sup> O. Möhler<sup>2</sup>

science, 2006



# Sulfate phase impact on cirrus formation

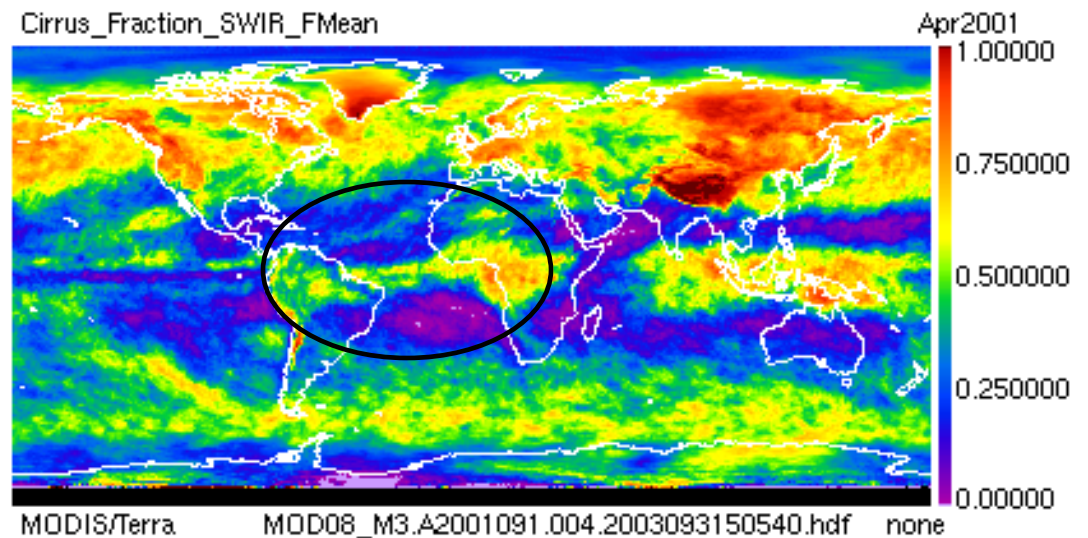
Regions where phase transition has the largest impact.  $\Delta M_{sd}/M$



Places having important phase transition all show larger cirrus cloud fraction.

The reverse is not necessarily true, because there are other factors

Cirrus cloud fraction from MODIS



# Summary

- Simulation of sulfate phase in  $\text{H}_2\text{SO}_4\text{-H}_2\text{O}$  system is developed. Inclusion of nitrate and organic aerosols is on the way.
- $\text{SDCF}_{\text{CRH}=0}$  is about 4% larger than  $\text{SDCF}_{\text{basecase}}$ ,  $\text{SDCF}_{\text{CRH}=\text{DRH}}$  is 6% smaller than  $\text{SDCF}_{\text{basecase}}$ .
- The percentage might vary if the model doesn't resolve the sulfate chemical composition, and hence hygroscopicity.
- Phase transition impact on SDCF has the important seasonal and regional variations, with larger effect (>20%) over south America, south Africa, east Asia, Europe and U.S. during summer time.
- The simulation results have important implications for understanding cirrus cloud formation.

# Acknowledgement

- NASA Atmospheric Composition Modeling and Analysis Program.
- NSF
- NOAA Climate and Global Change postdoctoral fellowship program under the administration of UCAR.

Assume same size distribution of dry particles, as RH changes from 0% to 80%, E increases by a factor of 2.7! For the same RH, chemical composition results in variation of E within 20%.

	RH=0				
	$m_r$	density ( $\text{gcm}^{-3}$ )	$r_{\text{eff}}$ ( $\mu\text{m}$ )	E $\text{m}^2(\text{g dry particle})^{-1}$	E $\text{m}^2(\text{gSO}_4^{2-})^{-1}$
AS	1.53	1.76	0.17	3.85	5.31
AHS	1.47	1.78	0.17	3.29	3.95
LET	1.51	1.83	0.17	3.53	4.55
SA	1.84	1.84	0.17	3.42	3.48
Water	1.33	1.00			

3.52 $\pm$ 0.24,  $\pm 7\%$       4.32 $\pm$ 0.79,  $\pm 18\%$

	RH=80%				
AS	1.41	1.30	0.24	9.62	13.28
AHS	1.38	1.30	0.25	9.84	11.81
LET	1.40	1.32	0.25	9.85	12.71
SA	1.37	1.24	0.30	15.49	15.80

11.2 $\pm$ 2.86,  $\pm 26\%$       13.4 $\pm$ 1.7,  $\pm 13\%$

In GOES-CHEM, E ~ 11.5, assume  $r_g=0.05$ , this study 0.07

Literature 8.0 - 16.0